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Size matters: biochemical mineralization and microbial incorporation of dicarboxylic acids in soil

Hamed Kashi · Sebastian Loeppmann · Jennifer Herschbach · Carina Schink · Wolfgang Imhof · Reza Mohsenian Kouchaksaraee · Michaela A. Dippold · Sandra Spielvogel

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Abstract The transformation and turnover time of medium- to long-chain dicarboxylic acids (DCA) in soil is regulated by microbial uptake and mineralization. However, the chain length of n-alkyl lipids may have a remarkable influence on its microbial utilization and mineralization and therefore on the formation of stable soil organic carbon from e.g. leave- needle- and root-derived organic matter during decomposition. To investigate their size dependent mineralization and microbial incorporation, four DCA of different chain lengths (12–30 carbon atoms), that were ¹³C labeled at each of their

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H. Kashi (⊠) · S. Loeppmann · J. Herschbach ·
R. M. Kouchaksaraee · S. Spielvogel
Institute for Plant Nutrition and Soil Science, Christian-Albrechts-University Kiel, Hermann-Rodewald-Str. 2,
24118 Kiel, Germany
e-mail: h.kashi@soils.uni-kiel.de

C. Schink · W. Imhof Institute for Integrated Natural Sciences, Department of Chemistry, University Koblenz-Landau, Universitätsstr. 1, 56070 Koblenz, Germany

M. A. Dippold

Biogeochemistry of Agroecosystems, Georg-August-University Göttingen, Büsgenweg 2, 37077 Göttingen, Germany

M. A. Dippold

Geo-Biosphere Interactions, University of Tübingen, Schnarrenbergstraße 94-96, 72076 Tübingen, Germany terminal carboxylic groups, were applied to the Ah horizon of a Fluvic Gleysol. Incorporation of ¹³C into CO₂ and in distinct microbial groups classified by phospholipid fatty acid (PLFA) analysis was investigated. Mineralization of DCA and incorporation into PLFA decreased with increasing chain length, and the mineralization rate was highest during the first days of incubation. Half-life time of DCA carbon in soil increased from 7.6 days for C12 DCA to 86.6 days for C₁₈ DCA and decreased again to 46.2 days for C22 DCA, whereas C30 DCA had the longest half-life time. Rapid and efficient uptake of C12 DCA as an intact molecule was observable. Gram-negative bacteria incorporated higher amounts of DCA-derived ¹³C compared to other microbial groups, especially compared to actinomycetes and fungi during the first phase of incubation. However, the incorporation of C₁₂ DCA derived ¹³C into the PLFA of actinomycetes, and fungi increased steadily during the entire incubation time, suggesting that those groups take up the ¹³C label from necromass of bacteria that used the C₁₂ DCA for formation of their lipids before.

Keywords ¹³C-labeling · Dicarboxylic acids · Phospholipid fatty acids · Microbial degradation

Introduction

An increasing number of studies emphasizes the significance of root-derived organic matter inputs as precursor material for soil organic carbon (Altmann et al. 2021a). One distinct class of compounds present in root-litter are long-chain dicarboxylic acids (DCA). DCA are known to be a major moiety (up to 50%) of suberin (van Bergen et al. 1998; Watanabe et al. 2013), a polyester which accumulates at the hypodermal/exodermal cell layers of the roots and that can make up to $62 \ \mu g \ cm^{-1}$ of a mature root, whereas lignin contributes in some cases only less than 50 µg cm^{-1} , depending on the plant species (Watanabe et al. 2013). Despite their quantitative importance in suberin, the fate of medium- to long-chain DCA in soil has received little attention, whereas the degradation of lignin and lignin monomers, as well as their contribution to SOM has been studied extensively (e.g. Zimmermann 1990; Bahri et al. 2006; Datta et al. 2017). Moreover, suberin-derived DCA have been used for decades as root-specific biomarkers to identify and quantify root inputs to sediments and soils (e.g. Nierop 1998; Rasse et al. 2005; Mendez-Millan et al. 2011; Altmann et al. 2021a). However, their use as root-biomarkers in soil requires appropriate knowledge about the fate of individual indicator molecules during degradation, which is still not well understood (Feng et al. 2010; Altmann et al. 2021a, b) determined the change of concentrations of rootand leaf-derived ester-bound lipids within one year of leave, needle and root degradation and found no linear trend, however, they focused solely on the depolymerization of the polymers cutin and suberin but did not study the degradation of individual monomers. The further cycling and turnover of DCA in soil, once the ester-bounds are cleaved is still not fully understood, which would be essential to their utilization as biomarkers for root-derived organic matter input and turnover.

The cleavage of ester-bonds and therefore release of cutin and suberin derived DCA by microbial cutinases and suberinases during litter decomposition is one possible source of DCA in the detritus sphere (Kolattukudy 1980). The composition of lipid monomers (e.g. ω -hydroxy fatty acids, dihydroxy fatty acids and DCA) released by cutinase CcCut1 showed a similar pattern to the monomers released by alkaline hydrolysis (Järvinen et al. 2009; Kontkanen et al. 2009). However, secondary formation via enzymatic oxidation of precursor compounds could also be a major, yet underestimated, source of DCA in soil (Ambles et al. 1994; Coon 2005; Rojo 2009). Microbial residues consist of long-chain n-alkanes, n-alkanols, and n-alkanoic acids. In the case of n-alkanes aerobic degradation usually starts with the oxidation of a terminal methyl group to render primary alcohols, which are further oxidized to the corresponding aldehydes, and finally converted into fatty acids (Rojo 2009). Fatty acids are conjugated to coenzyme A (CoA) and further processed by β -oxidation to generate acetyl-CoA (Watkinson and Morgan 1990; Ashraf et al. 1994; van Hamme et al. 2003; Wentzel et al. 2007). Both ends of the alkane molecule may be oxidized through ω -hydroxylation of fatty acids at the terminal methyl group, rendering an ω -hydroxy fatty acid. The latter is further converted into DCA which can subsequently be processed by β -oxidation (Watkinson and Morgan 1990; Coon 2005). This synthesis of medium- to long-chain DCA and cutin- or suberin-derived monomers by microorganisms has been described for pure cultures of some yeasts (e.g. Candida tropicalis or Yarrowia lipolytica), filamentous fungi and some bacteria present in soils and litter layers (Mauersberger et al. 1992; Eschenfeldt et al. 2003; van den Brink et al. 1998; Scott 2008; Van Bogaert et al. 2011). In the course of three enzymatic reactions, fatty acids are oxidized to DCA (Mauersberger et al. 1992; Eschenfeldt et al. 2003).

Additionally, microorganisms also possess the biochemical degradation of DCA in soil via β -oxidation, a process similarly demonstrated in pure cultures for preventing e.g. the accumulation of DCA in naturally occurring yeasts (Mauersberger et al. 1992; Eschenfeldt et al. 2003). The degradation of DCA occurred slower for long-chain DCA (C18, C20, C22), which showed a higher resistance to microbial degradation than those with short and moderate chain lengths (C_{2-12}) (Olivella et al. 2015; Saxena and Hildemann 1996). Degradation time may be expressed as halflife, defining the time required to degrade half the amount of the specific compound, which was applied and released as CO₂ (Jones and Shannon 1999; Boddy et al. 2007; Pearson and Eglinton 2000) estimated mean terrigenous residence times for higher plant wax (C24-33) derived long-chain n-alkanes and DCA from sediments to be less than 100 years. Average terrestrial residence times of vascular plant biomarkers deduced from n-C₂₈₋₃₀ alkanes and from stations directly in front of the river mouths range from 900 to 4400 years (Kusch et al. 2010).

Bacteria (According to PLFA analysis, especially gram-negative bacteria) were highly competitive for the uptake of some organic acids over other microbial groups (Gunina et al. 2014; Garcia-Pausas and Paterson 2011; Waldrop and Firestone 2004). The contribution of fungi to carboxylic acid-C utilization was less than that of bacteria, due to their lower competitiveness for water soluble substances (Gunina et al. 2014; Waldrop and Firestone 2004). To unravel the fate and transformation/incorporation of DCA-C by specific microbial groups, we established an incubation experiment using position-specifically labeled $D^{13}CA$. Position-specific labeling is a tool that has recently reached an increasing consideration in soil science (Fischer and Kuzyakov 2010; Dijkstra et al. 2011; Apostel et al. 2013; Dippold and Kuzyakov 2013). It overcomes the limitations of uniform labeling because it allows differentiation between the incorporation of molecule fragments vs. the incorporation of entire intact molecules (Apostel et al. 2015). PLFA analysis allows the reconstruction of the microbial community composition (Zelles et al. 1995; Zelles 1999), but in combination with ¹³C labelingi.e. ¹³C-PLFA analysis – it enables the tracing of substrate incorporation into microbial cell membranes and therefore the reconstruction of individual microbial groups (Glaser 2005). We selected four carboxylic group enriched fatty diacids as monomers, which we expected to be quite abundant in soils since they have either been derived from microbial synthesis or suberin decomposition (Banfield et al. 2018). The incorporation of the terminal enrichment into hydrolysable lipids like PLFA will reveal that at one of the two carboxylic groups is still present in this lipid, most likely involved in the ester bond (Deamer 2017).

We hypothesized that (1) DCA half-life time increases with chain-length of DCA due to decreased microbial accessibility, (2) bacterial groups are most competitive in the uptake of the DCA with lower chain length, but this uptake competitive advantage decreases with increasing DCA chain length, and (3) with increasing turnover time of DCA, the incorporation of DCA-C into actinomycetes would increase as they are known as a highly competitive group for microbial necromass utilization.

Materials and methods

Sampling site

Soil samples were collected from the Ah horizon of a Fluvic Gleysol located in the Middle-Elbe flood plain used as grassland at the end of October 2016 from 0–10 cm depth. The geographical location of the site was 12°75' northern latitude; 51°83' eastern longitude and 107 m above sea level, respectively (Gallin, Germany). The average annual temperature has been recorded as 8 °C (Weather station Zarrentin am Schaalsee). The collected soil samples were sieved to <2 mm, and stored for one day at +5 °C, except for subsamples used to determine the soil water content (dried at 105 °C following standard procedures). Physico-chemical properties of the soil sample are given in Table 1.

Experimental setup

Four position specifically ¹³ C labeled isotopomers of DCA with different chain lengths were used as tracer compounds, and two non-labeled DCA served as natural background substances. A total of 72 microcosms (6 DCA treatments × 4 replicates × 3 sampling dates) were prepared. The DCA applied to the microcosms were dodecane-dioic acid (${}^{13}C_{12}H_{22}O_4$), octadecane-dioic acid (${}^{13}C_2C_{16}H_{22}O_4$), docosane-dioic acid (${}^{13}C_2C_{28}H_{22}O_4$). The rationale behind our approach to use DCAs with similar chemical structure (homologous series) but increasing chain length is that the

Table 1Basiccharacteristics of soilsample of the experimentsite with total organiccarbon (TOC), cationexchange capacity (CEC)

Sand	Silt	Clay	рН	TOC	Total N
(%)	(%)	(%)		(%)	(%)
15	48	37	5.05	4.9	0.46
CaCO ₃	CEC	Na	K	Ca	Mg
(%)	(cmol _c kg ⁻¹)				
< 0.8	20.3	0.9	0.3	14.4	2.5

mean nominal oxidation state of the C atoms (average NOSC) in the respective molecule decreases with increasing chain length (Chakrawal et al. 2020). Hence, the average NOSC of the applied DCAs decreases in the order C_{12} DCA, C_{18} DCA, C_{22} DCA to C_{30} DCA. Molecules containing carbon atoms that are more reduced on average (i.e. have a lower average NOSC) tend to have higher oxidation activation energies (i.e. higher Gibbs energies for the oxidation half reaction). As enzymes catalyze reactions by decreasing the energy of activation, we assume that microorganisms incur higher metabolic costs when decomposing long-chain molecules from a homologous series compared to their shorter-chain representatives (Gunina et al. 2017; Chakrawal et al. 2020).

To produce the applied DCA, a novel synthetic protocol based on chain elongation of various precursors followed by a terminal ¹³C introduction using $K^{13}CN$ and subsequent hydrolysis to produce the respective DCA was designed (Schink et al. 2020).

Labeled DCA were weighed according to their molar mass to achieve an application rate of $100 \mu mol$ fatty acids to 1 kg soil and were mixed with 100 mg silt-sized quartz (Quarzwerke GmbH, Frechen, Germany).

Soil incubation was done in 50 ml microcosms. For this purpose, 20 g of dry soil with a mean bulk density of 1.5 gcm⁻³ was weighed and transferred to microcosms in four different layers (6 g, 6 g, 5 g, and 3 g). These four soil layers were separated by three layers of DCA compounds. The final acid layer was transferred with two times 0.5 ml ultrapure water, to achieve a weight of 20.1 g \pm 0.2 g. Soil water content was kept at 45% water holding capacity (WHC). The microcosms were incubated for 14 days at 8 °C, which is the mean annual temperature of the sampling site.

Additionally, microcosms in 100 ml crimped bottles were installed for ${}^{13}\text{CO}_2$ incubation. They received similarly the four layers of soil (totally 20 g), the interlayer of sand and DCA, and were adjusted to a soil water content of 45% (WHC).

 13 CO₂ sampling took place after 1, 2, 3, 4, 6, 8, 10, 15, 20, 25, 30, and 35 days. A gas chromatography syringe was used to transfer trapped CO₂ to 15 ml pre-evacuated exetainers at the same time for all four replicates. Thereafter, syn air flushing was done

for 45 s on crimped bottles, and then they were transferred to the fridge to be kept at 10 °C until the next sampling time. The destructive soil sampling was done after 3, 10, and 40 days and soils were frozen at -20 °C until PLFA analysis and freeze-dried just before extraction.

Quantification of soil mineralization

The CO_2 of the gas samples was quantified using GC-FID-ECD (Thermo Fisher Scientific, Bremen, Germany). In addition to the samples, four test gases with a known CO_2 concentration were measured as calibration gases. These external standards contained 400, 1500, 3000, and 4000 ppm CO₂. The starting temperature of the measurements was 50 °C. This was held for five minutes. The temperature was then increased to 130 °C at a rate of 100 °C per minute. The final temperature was held for two minutes for column reconditioning. Following the measurements, the CO₂ peaks of each chromatogram were integrated, and the areas determined using Agilent Chem-Station 10 (Agilent, Waldbronn, Germany). To quantify the CO_2 of the gas samples, a quadratic calibration curve was created using the four external standards of increasing CO₂ concentration. Finally, the amount of CO₂ contained in the microcosms was calculated using the quantified CO₂ production and the general gas law (Schink 2019).

Determination of ¹³C enrichment in respired CO₂

The determination of the respired δ^{13} C values was carried out at the center for stable isotope analysis (KOSI) at Georg-August University Göttingen. Briefly, 5.5 ml of the sampled gas was injected into a GC-C-IRMS equipped with an autosampler unit (AS 200) (Thermo Fisher Scientific, Bremen, Germany). We identified the peak areas and integrated these via Isodat 2.0, and referenced the measurements by a reference gas, which was injected before each sample application.

Calculation of ¹³C incorporation into CO₂

 13 C in CO₂ derived from the applied DCA was calculated according to a linear mixing model (Eqs. 1 and

2) (Gearing et al. 1991) where the CO_2 amount of the background sample in Eq. 1 was determined by Eq. 2.

$$\begin{bmatrix} C_{CO_2} \end{bmatrix}_{LAB} * r_{CO_2 - LAB} = \begin{bmatrix} C_{CO_2} \end{bmatrix}_{BG} * r_{CO_2 - BG} + \begin{bmatrix} C_{CO_2} \end{bmatrix}_{appDCA} * r_{appDCA}$$
(1)

$$\left[C_{CO_2}\right]_{LAB} = \left[C_{CO_2}\right]_{BG} + \left[C_{CO_2}\right]_{appDCA}$$
(2)

where $[C_{CO_2}]_{LAB;BG;appDCA}$ is the amount of C released in CO₂ of labelled samples, background samples, or derived from applied DCA, respectively (all in µg C g^{-1} soil). $r_{CO_2-LAB;CO_2BG;appDCA}$ is the at% ¹³C enrichment of the CO₂-eflux from labelled microcosms, background microcosms or derived from the applied DCA, respectively.

Double exponential kinetics was applied to calculate the corresponding estimation or $^{13}CO_2$ efflux rates from individual pools at time *t* (Eq. 3) (Kuzyakov 2011).

$${}_{x}{}^{13}CO_{2}(t) = {}_{x}{}^{13}C_{\max_a} * k_{a} * e^{-k_{a}*t} + {}_{x}{}^{13}C_{\max_b} * k_{b} * e^{-k_{b}*t}$$
(3)

where ${}^{13}CO_2(t)$ is the percent of ${}^{13}C$ incorporation into CO_2 at any time *t*; ${}^{13}Cmax_a/b$ is the maximal utilization of the fast and slowly mineralizable DCA pool, respectively; and k_a and k_b are the fast and slow decomposition rate, with

$$\lambda_{a/b} = \frac{ln2}{k_{a/b}} \tag{4}$$

where $\lambda_{a/b}$ is the mineralization half-time of the fast and the slowly mineralizable DCA pool (Gunina et al. 2017). The cumulative ¹³CO₂ efflux with time results as integral of Eq. 3.

PLFA extraction and preparation

PLFA extraction and purification was done according to Frostegard et al. (1991), modified by Gunina et al. (2014). Before extraction, 1 μ g per μ l solution of the first internal standard (phosphatidylcholinedinonadecanoic acid-IS1) was added to 6 g of each soil sample and extracted with a Bligh-and-Dyer extraction solution buffered by a citric acid buffer. After liquid-liquid and solid-phase purification steps, the fatty acids were saponified to free fatty acids and derivatized into fatty acid methyl esters (FAME) following the description of Knapp (1979). 15 μ L of the second internal standard (tridecanoate methyl ester -IS2, 1 μ g per μ l) and 185 μ L of toluene were added to the extracted FAME and transferred to 200 μ l inserts in 1.5 ml vials for gas chromatography (GC) analysis. In parallel to the samples, 5 external standards with 28 fatty acids were hydrolyzed, derivatized, and measured on both GC together with the samples.

PLFA identification

FAME-contents were identified and quantified via gas chromatograph-mass spectrometry (GC–MS) (GC 5890 with MS 5971 A, Agilent, Waldbronn, Germany) with a 30 m DB1-MS column, in the selected ion mode. FAME-contents were calculated by relating their areas to the area of IS2 followed by calibration via linear regression of the external standards. Extracted PLFA-contents were corrected based on the recovery of IS1. Agilent's MassHunter Software was used for integration.

Determination of ¹³ C values of PLFA

Measuring the FAME δ^{13} C analyzed by GC-C-IRMS consisting of an autosampler unit (AS 200) and combustion unit (Combustion Interface III) interfaced to a Delta plus isotope ratio mass spectrometer (all units from Thermo Fisher Scientific, Bremen, Germany). One µl sample were injected with 1 min splitless time into a liner heated to 280 °C. Peak separation was accomplished with two capillary columns (30 m DB-5MS followed by a 15 m DB-1MS, Agilent Waldbronn Germany) with helium (99.9996% pure) as carrier gas at a flow rate of 1.7 mL/min. Isodat 2.0 software was used for the calculation of δ^{13} C values (%) after the integration of FAME peaks. CO₂ reference gas with 99.995% purity was injected four times per chromatogram for a first calibration and drift correction.

Calculation of tracer ¹³C incorporation

The isotopic mixing model suggested by Gearing et al. (1991) was used to calculate the amount of

¹³C incorporation into PLFA (in % of applied tracer) (Eqs. 5 and 6).

$$\begin{bmatrix} C_{PLFA} \end{bmatrix}_{LAB} * r_{PLFA_LAB} = \begin{bmatrix} C_{PLFA} \end{bmatrix}_{BG} * r_{PLFA_BG} + \begin{bmatrix} C_{PLFA} \end{bmatrix}_{appDCA} * r_{appDCA}$$
(5)

$$\left[C_{PLFA}\right]_{LAB} = \left[C_{PLFA}\right]_{BG} + \left[C_{PLFA}\right]_{appDCA} \tag{6}$$

where $[C_{PLFA}]_{LAB;BG;appDCA}$ is the amount of C in PLFA of the labelled sample, or in the background sample or derived from the applied DCA (µg C per g⁻¹ soil), respectively. $r_{PLFA_LAB; PLFA_BG; appDCA}$ is the at% enrichment of the PLFA of the labelled microcosm or the background microcosm or derived from the applied DCA, respectively.

These equations quantify the incorporation of 13 C DCA into the respective fatty acids. Then the gram of 13 C incorporation has been divided by the amount of applied 13 C and multiplied with 100% to get the relative incorporation to the applied 13 C, also called tracer 13 C recovery in the respective fatty acid. This calculation is critical for the comparison between the individual DCA as it corrects the minimal differences in 13 C application of the individual tracer amendments. The incorporation is a quantitative measure of the total 13 C allocated into the pool of this membrane lipid classes.

We determined the relative incorporation into PLFA (F-value), reflecting a rather ecological parameter that describes the relative incorporation of the fatty acids of a distinct microbial group irrespective of its abundance, describing the replacement of ¹²C by ¹³C ($^{12}C/^{13}C_{repl}$) (Gunina et al. 2014) (Eq. 7).

$${}^{(12}C/{}^{13}C)_{repl} = \frac{\left[C_{PLFA}\right]_{appDCA}}{\left[C_{PLFA}\right]_{total}} * 100\%$$
(7)

where $({}_{X}{}^{12}C/{}_{X}{}^{13}C)_{repl}$ is the % of replacement of the unlabeled PLFA-pool by DCA-derived ¹³C-PLFA (µg ¹³C per soil); and $[C_{PLFA}]_{total}$ is the total amount of PLFA in the respective PLFA pool (µg C per soil).

Statistical analysis

A Nalimov outlier test at a 95% significance level was performed for mineralization and incorporation values (Lozán and Kausch 1998). The number of replicates has been considered before performing the outlier test. To disentangle the fate of the positionspecific labeled $D^{13}CA$ the respired ${}^{13}C-CO_2$, and incorporation into PLFA was investigated at different time points. The significance of differences was tested by repeated-measures analysis of variance (ANOVA) (Supplementary Table S1, S2). Tukey's post-hoc test was done after considering the assumptions of the ANOVA test (homogeneity and normal distribution) at 95% confidence level. The microbial grouping was conducted using factor analysis of PLFA contents. Microbial grouping was achieved if PLFA loaded on the same factor higher than 0.5 and if literature data on pure cultures revealed a common origin of the PLFA loading on the same factor (Apostel et al. 2013; Zelles 1999; Zelles et al. 1995). Statistica 12.0 software was used for all this statistical analysis.

Results

¹³CO₂ release from DCA mineralization

The mineralization of the terminal carboxylic groups of the DCA to CO_2 decreased with increasing chainlength of DCA for both, the cumulative (Fig. 1, top) and the rate of ¹³CO₂ (Fig. 1, bottom). The total ¹³CO₂ released during the incubation was considerably lower in C₁₈ DCA compared to the C₁₂ DCA (30-fold) and then further decreased for longer-chained DCA.

Mineralization of DCA to ${}^{13}\text{CO}_2$ was observed immediately after the application of compounds, and the rate of ${}^{13}\text{CO}_2$ released peaked within the first 24 h. The initial (maximal) respiration rate yielded 8, 0.12, 0.009 and 0.006% of applied tracer per day for C₁₂, C₁₈, C₂₂, and C₃₀ DCA, respectively. The DCA mineralization rate to ${}^{13}\text{CO}_2$ decreased in general rapidly after this initial peak, more intensively for C₁₂ DCA (320 times) compared to the other long-chain DCA (in average 9 times) (Fig. 1, bottom). In consequence, most of the ${}^{13}\text{CO}_2$ was respired during a short initial period.

Half-life time of DCA varied strongly between the two mineralization phases. During the fast mineralization, half-life time (λ_a) of DCA ranged between 1.8 and 2.3 d for C₁₂ and C₂₂ DCA, respectively, indicating no clear chain length effect for the first phase of DCA mineralization. The half-life time during the



Fig. 1 a cumulative respired ¹³CO₂ production in relation to the applied tracer (in % of applied DCA-¹³C) and **b** mineralization rate (proportion of respired ¹³CO₂ in relation to the tracer

applied in % of applied DCA/per day). Symbols represent mean values (n=4) and bars represent standard errors

Concorporation of the utilizable C pool, while ka and kb are DCA, respectively									
DCA compounds	C _{max} <i>a</i> (%)	$C_{max} b (\%)$	Ka (1/d)	Kb (1/d)	λ <i>a</i> (d)	λ <i>b</i> (d)			
C ₁₂	29.74	6.63	0.38	0.09	1.8	7.62			
C ₁₈	0.46	2.04	0.33	0.008	2.09	86.64			
C ₂₂	0.03	0.18	0.30	0.015	2.3	46.21			
C ₃₀	0.02	2.82	0.41	0.0004	1.67	1732.87			

Table 2 Decomposition of position-specific labeled $D^{13}CA$ over time using double kinetic coefficients. $C_{max} a$ and b are C incorporation of the utilizable C pool, while ka and kb are

the fast and slow decomposition rates. λa and b values represent the half-life time of the fast and slow decomposition of the DCA, respectively

slow (second) mineralization phase (λ_b) was noticeably longer compared to that of the fast (first) one (in average 239-fold). The half-life time of the second mineralization time varied widely between 7.62 and 1732.87 d for C₁₂ and C₃₀ DCA, respectively, and at least showed from shortest to longest DCA a pronounced effect of the chain-length (Table 2).

Incorporation of ¹³C labeled DCA into PLFA

The incorporation of ¹³C into the sum of all PLFA was highest for C₁₂ DCA (p<0.05) and had its maximum at the first harvesting time point (0.14% of applied tracer), thereafter decreasing significantly with time (Fig. 2). Overall, the highest ¹³C incorporation into PLFA was found in shorter-chain DCA compounds and always within the first few days of incubation. However, with increasing chain length the effect of time on the ¹³C incorporation disappeared. C₁₈ DCA was incorporated 12-fold, C₂₂ 45-fold and

 C_{30} 32-fold lower than C_{12} DCA into the sum of all PLFA, respectively (Fig. 2).

By combining the PCA based grouping of the fatty acids results with fatty acid fingerprints of pure-cultures we classified six microbial groups consisting of four groups of bacteria, including gram-positive 1 (G+1), gram-positive 2 (G+2), gram-negative 1 (G-1), and gram-negative 2 (G-2); one group of actinobacteria (Act), and one group of arbuscular mycorrhizal fungi (AMF) (Supplementary Table S3).

¹³C incorporation into PLFA for gram-negative bacteria was on average 6.9, 8.46 and 2.75 times higher than that of the gram-positive 1, gram positive 2 and gram negative 2 groups, respectively (Fig. 3). Overall, the incorporation of ¹³C₁₂ into PLFA of all bacterial groups was 21 times higher than ¹³C₁₈. Moreover, we determined a significant decrease of ~170% of ¹³C₁₂ DCA-derived C incorporation into PLFA of all bacterial groups from day 3 to day 40 (Fig. 3). The decrease in incorporation with







Labeled DCA tracers

Fig. 3 DCA-derived ¹³C incorporation into PLFA (in % of applied tracer) (mean \pm SEM) of different bacterial groups at different sampling times. Lower-case letters above error bars

indicate significant differences between the treatments and time of individual microbial groups at a level of p < 0.05, n = 4





Fig. 4 ¹³C incorporation into PLFA (in % of applied tracer) (mean \pm SEM) of actinobacteria (Act) and fungi/arbuscular mycorrhizal fungi microbial groups; data from position-specific labeled DCA at different time points. Lower-case letters

above error bars indicate significant differences between the treatments and time of individual microbial groups at a level of p < 0.05, n = 4

increasing incubation time was also visible for most higher chain length but in many cases not significant.

¹³C incorporation into the filamentous microorganisms changed only for ¹³C₁₂ DCA significantly over time, but, in contrast to the incorporation into single cell prokaryotes, increased over the duration of the experiment. Uptake of ¹³C₁₂ DCA into PLFA doubled for actinobacteria and arbuscular mycorrhizal fungi between day 10 and 40 (Fig. 4). Actinobacteria are the only group, where C_{12} and C_{18} -derived ¹³C was incorporated in similar amounts however with contrasting temporal pattern: whereas C₁₂-derived ¹³C incorporation increased with time, that of C22 DCA decreased with time resembling more the temporal pattern observed for free living prokaryotes. Filamentous microbial groups seemed to avoid the use of C_{22} DCA, which was lower than that of C_{30} DCA. Relative incorporation of ¹³C from DCA into PLFA (the isotopic F-value=fraction of tracer-derived ¹³C from total C of a pool) showed similar patterns for all microbial groups (supplementary Figure S1). Solely the filamentous actinobacteria did not show a decrease of relative incorporation of ¹³C from DCA into PLFA from the short-chain C12 DCA to the longer chain ones.

Discussion

Mineralization and total utilization

Most of the DCA mineralization was observed during the first days of the experiment leading to remarkably higher decomposition constants for the rapid phase of DCA mineralization. This suggests that the mineralization controlling factors are very different between the first and the second phase of decomposition. One often described phenomenon when adding substances to soils is the fact that mechanisms reducing microbial access to a potential C source are temporally slower than microbial uptake and utilization (Fischer and Kuzyakov 2010; Dippold and Kuzyakov 2013). This implies that during the initial phase mainly microbial uptake and metabolic potential to mineralize the DCA might have been the controlling factor for their decomposition, whereas their reduced access by e.g. hydrophobic interaction with hydrophobic parts of the soil organic matter might have added up on that increasingly with increasing incubation time leading to the reduced decomposition kinetics as observed in the second phase of the experiment (Kuzyakov 2011).

Mineralization of DCA displays a clear chain length effect, with cumulative mineralization over 40 days and the long-term decomposition constant K_b decreasing remarkably with increasing chain length. Compounds with a high hydrophobicity and thus a low water solubility are less susceptible to decomposition in soils (Miao et al. 2017). It was discovered that water-insoluble material reduces microbial activity as well as the rate of decomposition, particularly in organic soils (Spaccini et al. 2002; Miao et al. 2017). This suggests that hydrophobicity of added substances not only affects their own mineralization but also the mineralization of the adjacent SOC. So, decreasing mineralization of the DCA compounds with increasing number of carbon atoms can be explained by the increasing hydrophobicity with increasing chain length. The average NOSC is likely the other important factor in mineralization of DCA. Increasing the chain length in DCA compounds leads to the longer chains of $-(CH_2)$ - bonded carbon atoms that decreases the average NOSC of the overall molecule and consequently decreases its mineralization (Gunina et al. 2017). Nevertheless, both factors would suggest a continued decrease/increase with chain length but instead the chain-length effect was highly pronounced comparing the very long chain C_{30} DCA with the short-chain C₁₂ DCA but was not visible any more or not consistent when comparing the two mid-chain DCA C_{18} and C_{22} . This suggests that besides mere biophysical (hydrophobicity) and chemical (oxidation state) parameters, also clear adaptive mechanisms of microbial decomposition metabolism to defined chain length seem to exist, which partially overlay the availability mechanisms.

Several mechanisms are involved in cellular uptake of hydrophobic substrates in different microorganisms, including substrate solubilization by extracellular emulsifying agents (Lee et al. 2018; Atshaves et al. 2010), cell surface alteration to promote adhesion (Beopoulos et al. 2009), passive diffusion through the cell membrane (Huf et al. 2011), and transportermediated uptake (Thevenieau et al. 2007). However, it appears that the degradation pathways and transporters for middle-chain and long-chain fatty acids are distinct (Schaffer 2002; Werner and Zibek 2017).

Long-chain fatty acids are thought to enter cells through a mechanism that includes adsorption to the membrane, protonation, membrane flip-flopping, and desorption into the cytosolic space (Hamilton and Kamp 1999). Long-chain fatty acid uptake through cellular membranes is linked to esterification by acyl-CoA synthetases, according to studies in both prokaryotic and eukaryotic organisms (Schaffer 2002). Long-chain fatty acids are activated to acyl-CoA and must be transported to the mitochondrial interior via the carnitine shuttle in eukaryotic cells. In contrast, the common feature of short-chain and middle-chain fatty acids is their carnitine-independent uptake and, in case of eukaryotes, intramitochondrial activation to acyl-CoA thioesters (Schönfeld and Wojtczak 2016). Furthermore, unlike long-chain fatty acids, shortchain and middle-chain fatty acids' cellular metabolism is less dependent on fatty acid binding proteins (Schönfeld and Wojtczak 2016). Some evidences emerged that proteins are important mediators and/or regulators of trafficking of long-chain fatty acids into and within cells (Schaffer 2002; Atshaves et al. 2010). As a result, uptake of long-chain fatty acids needs more energy than middle-chain fatty acids (Schönfeld and Wojtczak 2016) – a potentially highly relevant factor explaining the reduced microbial utilization of the very long chain C_{30} DCA. We suppose that the lack of a chain-length effect in the range of our mid-chain DCA i.e. the observation that the increased hydrophobicity of C₂₂ DCA compared to C₁₈ DCA does not go along with a decreased decomposition, is linked to an alteration of chain-length specific uptake, intracellular activation or decomposition mechanisms not yet described in detail.

Utilization of DCA by the soil microbial community

Microorganisms, by producing microbial biomass and necromass have significant influence on SOM formation and cycling (Powlson et al. 2001; Koyama et al. 2018). Organic acids like DCA can be taken up by soil microorganisms whose residues in turn can be sorbed to soil particles or form biofilms on aggregate and particle surfaces, one key mechanism of SOM stabilization (Jones et al. 2003). The hydrophobicity of the cell surface is very important in the attachment to, or detachment from the surfaces (Wu et al. 2012). High cell surface hydrophobicity enables microorganisms to attach to hydrocarbon droplets on the surface or cells (Krasowska and Sigler 2014). Gram-negative bacteria increase cell surface hydrophobicity and their ability to form biofilms by releasing outer membrane vesicles (Baumgarten et al. 2012) which is most likely one of the key mechanisms explaining their high ability to use the DCA carbon for their metabolism. In contrast, Gram Positive bacteria's outer membrane is comprised of a peptidoglycan layer (Schaffer 2002), which is rather hydrophilic and has no general adaptation improving access to hydrophobic substances.

Fungi/arbuscular mycorrhizal fungi, as the eukaryotic group, was unable to take up similar amounts of the applied DCA compared to the prokaryotic groups in our experiment. Fungi are considered to be more hydrophobic than bacteria due to their hydrophobic rodlet layer (Kwan et al. 2006) and effective production of hydrophobins (Linder et al. 2005; Bayry et al. 2012). Therefore, microorganisms with hydrophobic surface preferentially interact with hydrophobic compounds, and likewise for hydrophilic microorganisms and compounds (Katsikogianni and Missirlis 2004). Regardless of molecular properties of an added substrate, the predominant bacterial biomass in the microbial community of our soil leads generally to a higher overall contribution of bacterial substrate use (Qiu et al. 2017). Additionally, eukaryotes also have a slower metabolism than that of prokaryotes because of their larger cell-size and higher cellular complexity (Bååth 1998; Rousk and Bååth 2007). The larger size of the eukaryotes' cell leads to a smaller surface to volume ratio (Apostel et al. 2013), and consequently, the uptake and incorporation of monomeric substrates into eukaryotic cell components takes longer (Moore et al. 2005; Reischke et al. 2014) found that loadings of small amounts of complex substrates did not support fungal growth, which may explain the low activity of fungal biomass in DCA utilization in this study. Bacteria rapidly take up labile organic compounds, whereas fungi are often associated with recalcitrant forms (Brant et al. 2006; Paterson et al. 2008; Gunina et al. 2017), a very general statement which we cannot confirm here for the case of hydrophobic long-chain DCA, which were still taken up to a higher degree by prokaryotes than eukaryotes in our study.

The lower uptake of applied DCA into PLFA by actinobacteria compared to other Gram positives indicated the preference of osmotrophic groups living at least partially as single cells in the soil solution for highly water-soluble, simple substrates (Gunina et al. 2014). Furthermore, actinobacteria and proteobacteria with the *alkB* gene were shown to be able to grow

on n-alkanes with lengths ranging from C_{10} to C_{19} as their sole source of carbon and energy (Ratajczak et al. 1998; Gibu et al. 2019). This can explain the higher uptake of C₁₂ and C₁₈ DCA in actinobacteria than in fungi/AMF. Furthermore, a second specific pattern was found for this group: Incorporation of ¹³C of C12 DCA into actinobacteria's PLFA increased significantly over the course of the experiment. This can be explained by the uptake of microbial necromass as a nutritional source of organics (Apostel et al. 2018). The high and rapid consumption of C_{12} DCA by the osmotrophic living gram-positive and gram-negative accompanied with a rapid turnover of their biomass results in an accumulation of considerable amounts of ¹³C-enriched necromass during the 40 days of the experiment. Actinobacteria, by producing lytic enzymes, have a wider metabolic ability than other bacteria in the uptake and use of necromass C (Ryckeboer et al. 2003). The increasing incorporation of DCA-C into actinobacteria PLFA in the second phase of the experiment demonstrates the beginning of this necromass recycling process.

Overall, we could demonstrate distinct mechanisms of DCA turnover in soils, which are controlled by chain-length and thus hydrophobicity at the one hand, but also by substrate-specific metabolic processes. These processes are likely driven by individual microbial groups which could induce divergent DCA pathways in soils. Moreover, our results indicate that the persistence of DCA in soil and therefore their applicability as biomarkers for root derived OC input, cycling and stabilization, might be a function of chain lengths. Altmann et al. (2021b) found a strong decrease of concentrations of cutin-derived dihydroxy fatty acids from Pinus needles after one year, that was less profound with increasing chain length. They explained this observation either by preferential degradation of the ester-bound compounds located on the surface of the cutin-polymers (Beaulieu et al. 2016; Altmann 2021b) or with the atomic oxygen-to-carbon ratio, which influences compound degradation rates (Kleber 2010; Altmann 2021b). Our results show that a chain-length dependent degradation occurs even if the cross-linking ester-bounds are already hydrolyzed and the fatty acids are available as monomers. This challenges the theory of a preferential degradation caused by location within the cutin- or suberin-polymer. However, general assumptions about the applicability of DCA as biomarkers and their stability in soil, solely based on incubation studies, have to be considered carefully and should be complemented e.g. with stable isotope approaches under field conditions. A schematic overview over these processes is illustrated in Fig. 5.

Conclusion

Position-specific labeling combined with ¹³C-PLFA and ¹³C–CO₂-analysis have revealed new information about the microbial transformation of DCA in soil. The biochemical degradation of DCA differing



in chain length demonstrated that C mineralization and cellular uptake of C_{12} DCA are higher than that of other long-chained DCA. The water-solubility and lower molecular weight but higher oxidation state of C_{12} DCA can be recognized as the most important factors accelerated its entry into metabolic cycles. The transformation pathway of long-chain DCA can differ from that of middle-chain DCA.

In most cases, mineralization to CO_2 and incorporation into PLFA decreased with increasing incubation time and a clear two-phase kinetic with altering factors controlling the DCA metabolism could be identified. Necromass utilization and lower general activity of actinomycetes and fungi/arbuscular mycorrhizal fungi changed this temporal pattern for the filamentous groups. Generally, the cellular structure and adaption mechanisms of the microbial groups played a crucial role in their ability to use the applied DCA, as reflected in the higher uptake of DCA into outer-membrane and thus hydrophobic gram-negative bacteria than into other microbial groups.

In conclusion, we found clear microbial preferences in the use of long chain DCA-derived C. Thus, microbial chain-length specific utilization preferences may significantly shape the DCA pattern observed within the free lipid fraction extracted from soils and the DCA composition after hydrolysis of cutin and suberin polymers. Hence, the specific degradation pattern of DCA in soil has to be considered when using them as biomarkers for root organic matter contribution to different SOM pools and should be complemented with stable isotope approaches.

Our results furthermore show that the longer-term stability of DCA derived carbon in soil is neither depending on the esterification of DCA in suberin-like structures nor on their cycling via the microbial carbon pump into microbial necromass. Instead it is rather a function of their chain-length dependent hydrophobicity and as such a recalcitrance based on their own aliphatic structure.

Author contributions All authors contributed to the study conception and design. Material preparation was performed by Jennifer Herschbach, Carina Schink and Wolfgang Imhof. Data collection and analysis were performed by Hamed Kashi, Michaela A. Dippold and Sandra Spielvogel. The first draft of the manuscript was written by Hamed Kashi and all authors commented on previous versions of the manuscript. All authors read and approved the final manuscript.

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Data availability The datasets generated during and/or analysed during the current study are available from the corresponding author on request but will be uploaded during the revision process to be available for readers when the publication gets online.

Declarations

Conflict of interest The authors have no relevant financial or non-financial interests to disclose.

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